

Suzanne A. Pierce, Shanna B. Evans, and John M. Sharp jr. Jackson School of Geosciences, The University of Texas at Austin Groundwater systems in the Cuatrociénegas basin: comparing two conceptual flow models and related implications for groundwater management strategies





# The Cuatrociénegas Aquifer System

What we know...

### Complex Groundwater Behavior

Karst features are adjacent to playa lakes within the valley and the actual groundwater flowpaths are unclear. Fracture dominated aquifer overlain by porous aquifer. Salinity issues. This system exhibits evidence of widely varied processes in a very small scale.

### Groundwater Dependent Ecosystem Setting

In the Chihuahuan Desert, over 167 Springs and surface water expressions unique level of biodiversity.

### Human Induced Change within the System

Naturally a closed or partially closed system, but due to use outside of the valley the system has become an artificially open one. Recently, decreases in water levels within the valley are disrupting the surface ecosystems and raising doubts about the future viability of continued irrigation within the valley.

## The Cuatrociénegas Basin Geologic Setting

![](_page_2_Figure_1.jpeg)

Steeply dipping anticline with alluvial valley fill in between

Bounding mountain ranges are dominated by **massive limestone** and **detrital shed** (McKee, 1990).

The **interior** of the valley is primarily unconsolidated **evaporite deposits**.

### **Boundaries**

Northwest = thrust faults along the flanks of the Sierra de San Madera (Lesser, 2001). Southern = combination of thrust and normal faulting from uplift of the Sierra de San Marcos or Sierra de la Fragua (Lesser, 2001).

# Hydrogeologic Inputs + Outputs

#### Recharge

Probably precipitation in upland areas Arid climate = Chihuahuan Desert Annual precipitation = 200 mm Temperature ranges = 44°C to 0°C in the winter (Minckley, 1969 and Pronatura, 1998).

#### Discharge

Natural = Evaporation and Evapotranspiration, Artificial = Irrigation and Municipal Demand (water table drop 30-60 cm past 5 years – Lesser, 2001)

![](_page_3_Picture_5.jpeg)

# Hydrogeologic Character

![](_page_4_Figure_1.jpeg)

Storage = 0.022 - 0.0003

But no good estimates for transmissivity evapotranspiration, or discharge rates

# Hydrostratigraphic Units

Tertiary Quatern ary	Alluvium & Colluvium	Uncolsolidated gravels, sands, evaporite deposits	Permeable
	Lake Deposits	Clays	Impermeable
Cretaceous	Eagle Ford	Sandy clays, lutites and limestone	Impermeable
	Buda	Limestone	Permeable
	Del Rio	Lutites	Impermeable
	Georgetown	Limestone	Permeable
	Kiamichi	Lutites & limestones	Impermeable
	Aurora	Platform limestones	Permeable
	Cañon de Barril Viejo (or La Pena)	Mixed terrigenous and marine carbonate (interfingered)	
	Cupido	Graded limestone cyclical lithofacies transitions, massive limestone	Permeable
Late Jurassic	La Virgen	Gypsum interbedded with limestone sabkha on alluvial surface	Semipermeable
	La Mula	Sand-shale facies and lime-shale facies detrital deposition	Impermeable
	Padilla	Dolomite lagoonal environment	Semipermeable
	San Marcos	Basal arkose sandstone	Semipermeable
Pre-Jurassic	Basement	Igneous lithologies primarily dacites & granidorites	
(Modified primarily from McKee, et. al., 1990 and Martinez, 2000)			

### The basin is a naturally undrained and closed

![](_page_6_Picture_1.jpeg)

![](_page_6_Figure_2.jpeg)

Under natural conditions the principle means of discharge within the basin would be through evaporation & evapotranspiration

Today it is behaving as an artificially opensystem with the largest component of discharge due to inter-basin transfers for use in irrigation and municipal water sources.

# Spatial evaluation of hydrochemical data

Flanks of mountains exhibit lower salinity waters and Water quality degrades as the central sections of the basin are approached

![](_page_7_Figure_2.jpeg)

Conductivity Distribution in µS

![](_page_7_Figure_4.jpeg)

Low High

Springs near flanks = lowest

Playa, karst features, basin floor = highest

### Hydrochemistry holds with a simple closed basin interpretation

![](_page_8_Figure_1.jpeg)

![](_page_9_Figure_0.jpeg)

Playa basin features forming adjacent to karst features does not fit with the undrained, closed basin flow regime

![](_page_10_Picture_0.jpeg)

Shallow groundwater flows within basin are dominated by subflow processes

![](_page_10_Figure_2.jpeg)

# **Doline Formation**

Diffuse recharge becomes concentrated in subcutaneous zone

![](_page_10_Picture_5.jpeg)

This supports an epi-karst aquifer overlying a deeper seated groundwater source

![](_page_11_Picture_0.jpeg)

# **Doline Direction**

![](_page_11_Figure_2.jpeg)

# ... Dissolution & Mixing

![](_page_11_Picture_4.jpeg)

![](_page_12_Picture_0.jpeg)

If solution mixing contributes to doline formation, then what is the secondary source of groundwater?

![](_page_13_Figure_0.jpeg)

### Human Impacts Irrigation canals discharge points and recharge points

Agriculture is probably the major demand for water from the basin, but the canals also provide distributed recharge and return flows

![](_page_14_Figure_2.jpeg)

![](_page_14_Picture_3.jpeg)

# In Addition, there are a wide-range of practical problems to be solved. For example,

![](_page_15_Picture_1.jpeg)

Water Quality Analysis for Ejido Antiguos Mineros

![](_page_15_Figure_3.jpeg)

Conductivity Levels in Norias of Antiguos Mineros

![](_page_15_Figure_5.jpeg)

# Accessible Water Resource Management

Tools can be developed using simple water quality parameters and spatially related datamodels

- 1. Develop a water budget/balance for the basin (focus on surface water flows and ET estimates to start)
- 2. Clarify southwestern boundary conditions (geophysical methods)
- 3. Use existing hydrochemical data to model flow paths
- 4. Quantify hydrologic parameters (Drill long-term groundwater monitoring wells and conduct pump tests)
- 5. Finally, priority tasks should be internal to the CC basin, the effect of drawdown due to extra-basin pumping is most likely a secondary issue, for now.